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全球气候变化下土壤呼吸对温度和水分变化的响应特征综述*

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摘要: 气候变暖、降水格局变化等是气候变化的主要表现形式, 也是影响土壤呼吸主要的非生物因素, 探明气象条件(温度、水分)对土壤呼吸影响及作用机制是理解陆地生态系统碳循环的重要内容之一。本文对近年来国内外学者关于温度和水分对土壤呼吸的影响及机制的研究进展进行综述。结果表明, (1) 气候变暖与土壤呼吸存在正反馈调节, 但温度适应性削弱了二者的反馈关系。增温时长和土壤碳储量不同导致温度对土壤呼吸的影响具有时空差异。土壤呼吸对温度适应性机制主要包括土壤微生物适应性、底物消耗和土壤矿物质活化等。(2) 降水对土壤呼吸的作用取决于土壤初始水分含量。当土壤含水量低于萎蔫系数时, 降水不仅增加土壤含水量还可促进土壤呼吸, 在土壤含水量接近田间持水量时土壤呼吸达到最大值, 当土壤含水量达到饱和值时土壤呼吸又会受到抑制。水分对土壤呼吸影响机制主要为替代效应与阻滞效应、底物供给、微生物胁迫以及根系响应等。(3) 土壤呼吸与土壤温度、水分的耦合关系取决于土壤水热因子配比, 当土壤温度成为胁迫因子时, 降水引发的土壤水分含量升高对土壤呼吸的激发效应被低温的负面影响所抑制; 当土壤水分成为胁迫因子时, 气候变暖引发的土壤温度升高对土壤呼吸的促进作用被干旱的负面影响所抵消, 进行土壤呼吸研究时需充分考虑土壤温度和水分的交互作用。为更全面深入地明晰陆地生态系统土壤碳排放扰动因素, 未来气候变化下土壤呼吸与环境关系等相关领域研究应为重点方向, 一是加强多因素交互作用对土壤呼吸影响的研究, 并定量化研究土壤呼吸组分; 二是持续关注土壤呼吸对土壤初始温度和温度波动的响应特征, 探索生物多样性或群落结构组成对土壤呼吸的影响。

关键词: 土壤呼吸; 温度; 降水; 响应特征

A Review of the Response Characteristics of Soil Respiration to Temperature and Moisture Changes under Global Climate Change

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Abstract: Warming of the climate and changes in precipitation patterns are major manifestations of climate change and abiotic factors affecting soil respiration. Authors presents a systematic analysis of recent research advances on the effects and mechanisms of temperature and moisture on soil respiration. The results show that: (1) there is positive feedback between soil respiration and climate warming, but the temperature adaptation weakens this positive feedback. The effect of temperature on soil respiration varies spatially and temporally due to the different duration of warming and soil carbon storage. The main mechanisms of soil respiration adaptation to temperature include soil microbial adaptation, substrate depletion and soil mineral activation. (2) The effect of precipitation on soil respiration

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depends on the initial soil water content. When soil water content is lower than the wilting factor, precipitation not only increases soil water content but also promotes soil respiration, reaching a maximum when soil water content is close to the field holding capacity, while soil respiration is inhibited when soil water content reaches saturation value. The main mechanisms by which water affects soil respiration are substitution and blocking effects, substrate supply, microbial stress and root response. (3) The coupling of soil respiration with soil temperature and moisture depends on the ratio of soil water and heat factors. When soil temperature becomes a stress factor, the stimulating effect of increasing soil water content induced by precipitation on soil respiration is suppressed by the negative effect of low temperature. When soil moisture becomes a stress factor, the promoting effect of increased soil temperature due to climate warming on soil respiration is counteracted by the negative impact of drought. The interaction between soil temperature and moisture should be fully considered when studying soil respiration. In order to understand the disturbance factors of soil carbon emissions in terrestrial ecosystems, this paper proposes that future research on the relationship between soil respiration and the environment under climate change. Firstly, strengthen the research on the effects of multi-factor interaction on soil respiration and quantify the soil respiration components. Secondly, continue to pay attention to the characteristics of soil respiration in response to initial soil temperature and temperature fluctuations, and to explore the effects of biodiversity or community structure composition on soil respiration.

Key words: Soil respiration; Temperature; Rainfall; Response characteristics

全球气候变化导致温度升高及降水时空格局变化，影响了土壤有机底物的分解转化和供应，进而对土壤呼吸产生重要影响^[1]。全球陆地生态系统年均土壤碳排放总量为 $78\sim100\text{Pg}\cdot\text{a}^{-1}$ ，占陆地生态系统呼吸总量的 65%~85%^[2]，是化石燃料燃烧产生 CO_2 的 10 余倍^[3]。土壤呼吸作为全球陆地碳循环最大的流动通量之一，其极微小的变化都可能对大气 CO_2 浓度产生重大影响，对全球碳收支产生反馈效应，进而影响气候变化^[4]。

土壤呼吸速率受多种环境因素影响，其中温度和水分是两个最重要的非生物因素^[5]。研究普遍认为土壤呼吸与气候变暖存在正向反馈调节，气候变暖引起土壤温度升高促进了土壤呼吸^[6]，且二者呈指数相关性^[7]；也有研究表明，土壤呼吸与土壤温度遵循单峰型高斯响应模型^[8]。此外，气候变暖同时全球降水格局也随之改变，土壤水分对土壤呼吸影响的研究也备受关注，但其研究结果不尽相同，Zhang 等^[9]基于华北平原草地生态系统的 8a 降水试验，发现土壤水分增加可显著促进土壤呼吸；而 Han 等^[10]研究发现滨海湿地降水提高土壤含水量且趋于饱和时，导致土壤环境缺氧，降低了土壤呼吸。气候变化对土壤呼吸的影响是多因素共同作用的结果，全面了解其对土壤呼吸的影响及机制作用，有助于理解气候变化与陆地生态系统碳循环的关系。

目前，土壤温度和水分变化对土壤呼吸的影响已有大量研究，但仍缺少较为全面系统的总结。本文基于近年来已发表的相关研究成果，综述全球气候变暖背景下温度和水分对土壤呼吸的影响及其作用机制，以期为未来土壤呼吸对气候变化响应研究和评估预测提供理论参考。

1 温度对土壤呼吸的影响

1.1 综合影响

全球气候变暖引起土壤温度升高促进了土壤呼吸，大气 CO_2 浓度上升导致的温室效应会进一步加剧气候的变暖^[4]，因此陆地生态系统土壤碳排放与气候存在正反馈调节^[6]。增温时长可改变温度对土壤呼吸的影响。Melillo 等^[11]基于中纬度阔叶林长期土壤增温试验，发现温度对土壤呼吸的影响因增温时间延长而有差异，10a 增温大量消耗土壤有机碳，土壤呼吸逐渐减弱但仍显著高于未增温处理，17a 增温限制土壤可用性底物，增温处理与未增温处理下土壤呼吸差异不显著。Dacal 等^[12]基于旱地生态系统 10a 增温试验，验证了短期增温（≤2a）引起土壤温度升高，导致土壤呼吸增加，但短期增温对土壤呼吸的影响未随增温时长的延长而持续提高。短期增温主要通过消耗土壤易分解碳和刺激微生物代谢来刺激土壤呼吸，随增温时间延长，土壤微生物生物量下降及碳分解基因转变是土壤呼吸响应差异的主要原

因^[13]。

土壤呼吸对温度升高的反馈因土壤碳储量丰富度而存在差异。Luo 等^[14]基于美国草原生态系统 2a 增温试验,发现刈割样地相比非刈割样地土壤碳输入量低,土壤微生物群落组成变化可引发土壤呼吸下降以应对有限的土壤底物供应,表现出土壤含碳量低的样地比含碳量高的样地更能适应变暖。此结论应用于更大的区域尺度中,即土壤温度较高的温暖地区相较土壤温度较低的寒冷地区更适应气候变暖,具有更低的土壤呼吸温度敏感性。区域土壤碳储量为生态系统光合增益与土壤呼吸损失的碳含量差值^[6],而温度是控制高纬度和高海拔地区土壤碳储量的主要因素^[6,15],土壤碳储量随环境变冷呈指数级增长^[16]。大量基于田间试验、空间观测和数据整合分析的研究表明,气候变暖引发土壤温度升高可造成寒冷气候区土壤高有机碳储量大量排放^[16-17]。

1.2 温度敏感性

土壤呼吸温度敏感指数(Q_{10})是指温度每升高10℃土壤呼吸增加的倍数^[18]。土壤呼吸由异养呼吸和自养呼吸两部分组成,异养呼吸对温度的敏感性高于自养呼吸对温度的敏感性^[19-20]。土壤异养呼吸的温度敏感性与土壤有机质分解的温度敏感性相对应,取决于土壤微生物对土壤有机质及植物凋落物的分解能力,受低温、土壤含水量、底物基质可用性和微生物群落结构的限制^[21];土壤自养呼吸的温度敏感性一般指植物根系呼吸对温度的敏感性^[22],取决于植物根系生物量及根际共生体的类型,主要依赖植物光合产物的向下运输^[23],土壤温度升高对生物和非生物因子的影响差异导致土壤呼吸组分对温度敏感性的差异^[24]。目前野外原位测量土壤呼吸组分的技术还很难精细区分二者,现阶段普遍采用土壤呼吸组分拆分法,但由于生物及土壤化学因素的影响该方法存在一定误差^[25-26]。

Q_{10} 具有较强的季节和年际变化规律^[8,27]。Chen 等^[27]基于青藏高原东部亚高山森林生态系统2004-2006年生长季内的研究,发现 Q_{10} 值存在季节和年际变化,且随土壤温度升高而减小。卢闯等^[8]连续监测北京夏玉米-冬小麦轮作田土壤呼吸也得出相同结论,冬小麦种植期内 Q_{10} 与土壤温度呈显著负相关关系。可用底物供给量、土壤温度和其他限制因素共同影响土壤呼吸通量,而 Q_{10} 与植被碳输入的数量呈正相关^[28]。在较冷季节,不论是农田生态

系统还是森林生态系统,植物的枯枝落叶大都汇集于土壤表层,一定程度上增加了土壤可用性底物供给量,土壤可用性底物随温度变化导致 Q_{10} 增加^[29]。在不同气候梯度生态系统的土壤中,冷生物群落(北极和北方)较暖生物群落(温带和热带)具有更高的土壤呼吸温度敏感性^[17-18],年平均温度为5~10℃的高纬度地区 Q_{10} 随年平均温度增加而增加,该地区深层有机质存储于土壤矿物中,因此,土壤可用性底物扩散缓慢且胞外酶活性低,加上土壤通气、透水性能较差,土壤有机底物碳密度更高^[22],丰富的土壤碳储量和递增的温度使生态系统对温度的响应增加^[30]。

1.3 温度适应性

土壤呼吸温度适应性指土壤适应外界温度变化的能力^[18],即随土壤温度升高或增温时间延长,削弱土壤呼吸与土壤温度变化的正反馈,二者呈非线性关系,土壤呼吸速率增长幅度逐渐减缓^[14,31],表现为适宜范围内增温促进土壤有机质分解速率, Q_{10} 降低,当增温幅度超过一定阈值后土壤有机质分解速率逐渐减弱甚至转为抑制作用^[32]。旱地生态系统 10a 增温试验显示,短期($\leq 2a$)土壤温度升高显著促进土壤碳排放,随增温时间延长至8~10a,增温对土壤呼吸的促进作用消失^[12]。付薇等^[33]基于模拟增温试验,发现连续 2a 增温使土壤呼吸对温度产生了一定适应性,增温对大豆生长季土壤呼吸无显著促进作用。

1.4 适应性机制

1.4.1 土壤微生物适应性

土壤温度升高诱导微生物群落组成和功能结构的适应性变化,是导致土壤呼吸温度敏感性降低的主要原因^[34-35]。依据生态代谢理论(MTE),温度是生物活性的主要驱动力,不同生态系统对土壤微生物群落影响不同^[35]。气候变暖导致的土壤温度升高作为确定性的环境指标,显著影响土壤微生物群落结构^[36],即长期增温导致微生物生物量降低且直接影响土壤微生物多样性^[37]。其次,增温引起的土壤水分下降会改变土壤微生物群落组成和结构,增加耐旱属真菌相对丰度^[38],削弱变暖对土壤呼吸的刺激效应;此外,土壤碳输入量因作物种类不同而不同^[39],变暖通过改变地上植物群落结构而影响地下微生物群落结构组成,具体表现为长期增温使样地植物群落中C3植物生物量显著降低,向C4植物转

变^[35]，植物群落变化引发的微生物群落功能结构改变，是造成土壤微生物热适应的主要原因^[35]。

1.4.2 底物消耗机制

土壤温度适应性取决于易分解底物的数量和质量^[40]。土壤低温可抑制土壤有机质分解转化，短期增温会解除这种抑制作用，增加土壤可利用性底物，促进土壤呼吸作用，随着增温时间延长，土壤可利用性底物减少，土壤呼吸速率及土壤呼吸温度敏感性随之降低，体现出更强的温度适应性^[40-41]。与北极高纬度寒冷地区相比，热带低纬度温暖地区因土壤含碳量低而表现出更强温度适应性^[14,18,42]。Eliasson 等利用生态系统 GDAY 模型模拟升温对土壤碳含量的动态影响，结果表明，当环境温度升高 5℃时，土壤呼吸速率在 1a 后增加了 60%，而 10a 后仅增加了 3%^[41]，说明持续变暖使土壤矿化速率加快，消耗了不稳定有机碳库，导致土壤异养呼吸减弱^[40]，体现了土壤呼吸对增温的适应性。

1.4.3 土壤矿物质活化机制

升温可显著降低土壤微生物生物量氮，通过提高土壤氮的矿化速率而增加土壤氮素含量^[43]。升温条件下，土壤微生物以提高氮利用率的方式维持自身生理代谢以适应温度变化，从而抑制了微生物氮的固定^[44]。一定温度范围内，氮素转化酶活性随温度的升高而增加^[45]，促进土壤氮矿化^[46]。土壤氮素含量增加使得土壤 C : N 降低，促进植物对氮素的吸收^[47]，减少了植物光合产物向地下分配的比例，使得植物根系呼吸受到抑制，另一方面，根系分泌物的减少也会显著降低微生物可用性底物量，抑制土壤微生物活性^[48]，进一步降低土壤异养呼吸。

2 水分对土壤呼吸的影响及其机制

2.1 水分对土壤呼吸的影响

气候变化加剧水文循环，导致全球或区域降水格局发生改变^[2]。降水作为陆地生态系统土壤水分的主要来源，不仅增加土壤湿度，也促进植物地上部分有机残体向地下部分运输，为土壤呼吸提供主要底物^[49]。随着水分下渗迁移，土壤水分含量短期内迅速增加，土壤通透性和土壤中可溶性有机质浓度等理化性质发生改变^[50]，对土壤呼吸产生显著影响^[51]。杜珊珊等^[52]基于黄土高原雨养农区不同耕作模式下降水对土壤呼吸的影响研究，发现降水抑制了土壤呼吸。但王兴等^[53]发现陕北黄土丘陵区自然撂荒 12a，降水量的增加显著促进土壤呼吸。因此水分对土壤

呼吸的影响可能因植被类型、降水强度以及历时长短不同而产生差异。

降水对土壤呼吸促进还是抑制取决于土壤初始水分含量^[54]，目前学术界普遍利用抛物线模型模拟呼吸速率对土壤水分的响应^[2,55]，主要分为 3 个阶段：(1) 土壤含水量低于萎蔫系数时，土壤孔隙水分不连续，底物扩散受到抑制，土壤处于干旱胁迫状态^[56]。随着土壤含水量增加，底物扩散抑制作用得到改善，底物有效性提高，促进土壤呼吸；(2) 土壤含水量增加但仍低于田间持水量，土壤水分不再抑制底物和 O₂ 扩散，微生物也不再受干旱胁迫制约，植物根系和微生物活性达到最强，土壤呼吸速率最大^[8]；(3) 土壤含水量进一步增加，土壤处于饱和或积水状态，O₂ 扩散受阻，土壤整体处于厌氧状态，根系呼吸和微生物活动受到极大抑制^[57]，土壤呼吸随土壤含水量增加逐渐减弱^[52](图 1)。与年平均降水量 >500mm 地区相比，年平均降水量 <500mm 的半干旱或干旱地区降水可显著促进土壤碳循环，增加土壤呼吸^[54]。海滨湿地土壤因受浅层地下水位影响其土壤含水量较高，降水后土壤迅速饱和，土壤呼吸与土壤含水量呈显著负相关^[58]。降水明显抑制湿润区土壤呼吸及其分量，使处于饱和或积水状态的土壤呼吸温度敏感性显著提高^[7]，但土壤呼吸温度敏感性不会随土壤含水量的持续增加而升高。总体上看，土壤含水量与 Q₁₀ 呈二次曲线关系，当土壤湿度超过土壤田间持水量，降水持续增加导致土壤含水量升高的同时会降低土壤温度，土壤呼吸受到抑制，Q₁₀ 值减小^[54]。

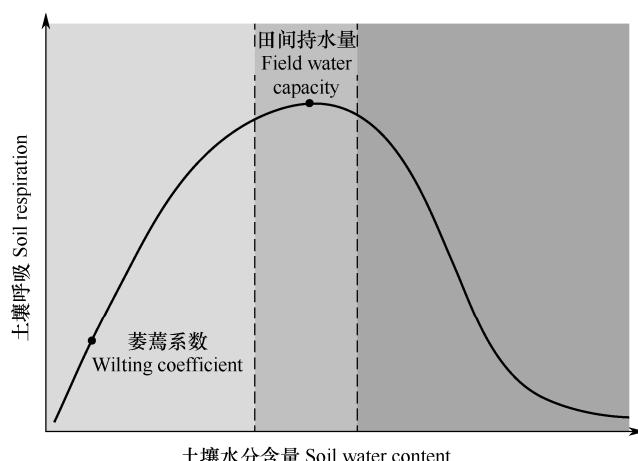


图 1 土壤呼吸对土壤水分变化的响应

Fig. 1 Soil respiration response on the changes of soil moisture

2.2 水分对土壤呼吸的影响机制

2.2.1 替代效应与阻滞效应

土壤处于干旱状态时, 孔隙中存在大量空气, 当降水下渗填充土壤孔隙后, 包括 CO₂ 在内的空气被替代并排出, 短时间内可激发土壤呼吸^[55]。但这并非真正意义的土壤呼吸而是物理排气过程, 是土壤孔隙对水分迅速增加的短暂响应, 此时土壤呼吸强弱取决于土壤干旱程度和理化性质^[59]。土壤湿润条件下, 降水导致土壤水分饱和使土壤整体处于厌氧状态, 潜在抑制土壤有效呼吸^[60]。此外, 高强度降水也可降低土壤通气性能, 通过阻碍 O₂ 在土壤中的扩散, 从而抑制土壤呼吸^[61]。

2.2.2 底物供给机制

土壤含水量增加使土壤团聚体因持续破裂而释放出较多有机碳是促进土壤呼吸的主要原因^[62]。研究发现, 干旱复水可使土壤微生物生物量和可溶性有机质分别增加 9.5% 和 12.3%^[63]。1958 年, Birch^[64]发现土壤干旱复水后有机碳矿化速率迅速加快, 土壤呼吸速率短时间达到峰值。土壤干旱状态时有机质大部分被土壤团聚体或胶体包裹, 土壤有机质被包裹的紧密程度与土壤含水量呈负相关关系, 土壤可用性底物降低将抑制土壤呼吸^[63], 此时增加土壤含水量使土壤团聚体结构破裂, 其比重(相对于土壤总重量)从 30% 减少到 21%^[65], 土壤微生物可利用性底物的有效面积增加, 土壤呼吸速率增加^[63]。但随着干湿交替次数增加, 土壤可用性底物持续减少将降低 CO₂ 排放量, 可很好地解释“birch 效应”中干旱复水刺激土壤 CO₂ 脉冲释放随时间减弱的变化趋势^[63]。

2.2.3 微生物胁迫机制

土壤微生物对干旱具有极强适应性, 可随时进入休眠或代谢缓慢阶段以适应干旱环境^[66]。休眠状态下土壤微生物活性极低, 干旱胁迫土壤中约 90% 的微生物处于休眠或非活性状态^[67], 此时微生物量减少且个体呼吸微弱, 土壤呼吸降低。降水后土壤水势短时间迅速上升, 部分土壤微生物通过破裂细胞释放胞内溶物的方式降低水势^[68], 而部分微生物则启动体内渗透调节机制应对土壤高水势, 其释放的大量含碳化合物则被余下的微生物分解利用^[69], 土壤呼吸升高。此外, 水分缺失导致的干旱胁迫可直接改变土壤微生物的群落组成和功能结构^[38], 土壤微生物优势类群由细菌转变为真菌, 并增加耐旱

型真菌的相对丰度^[70]。

2.2.4 根系响应机制

土壤孔隙是容纳水分和空气的空间, 同时也是植物根系活动的重要场所^[71]。土壤水分状况对土壤通透性有直接影响, 土壤水分含量与 O₂ 含量成反比, 而植物根系正常生长活动需要充足的 O₂ 供应, 当土壤水分饱和或处于积水状态时将抑制 O₂ 扩散, 因此长期的淹水状态会降低土壤根系呼吸。此外, 干旱胁迫时, 植物通过调控气孔闭合的方式降低蒸腾作用^[72], 此时植物光合作用受到抑制, 降低了光合产物向地下根系的分配能力, 植物根系生物量减少, 通过土壤根系呼吸受到限制的方式减弱植物代谢以应对干旱胁迫^[9,73]。

3 温度和水分交互作用对土壤呼吸的影响

在全球气候变化的过程中气温升高和降水格局变化是同时发生的^[74], 且气象条件(温度、水分)是影响土壤呼吸变化的主要环境因子^[2,75](图 2)。陆地生态系统土壤呼吸的变化并非单因素变化的作用结果, 而是各因素间交互作用或累积效应的共同体现, 对土壤呼吸造成协同或拮抗效应^[7-8]。

一般来说, 土壤呼吸随土壤温度升高呈指数增长, 随土壤水分增加呈抛物线变化^[2], 但温度对土壤呼吸的影响受限于土壤水分有效性^[35,76]。土壤干旱胁迫可通过降低土壤可溶性底物的扩散和胞外酶的活性, 以及抑制微生物和植物活性来降低土壤呼吸温度敏感性 (Q₁₀)^[55]。此外降水可部分抵消高温下土壤蒸散造成的水分损耗, 在补充土壤水分的同时增加土壤水分有效性, 二者交互作用可促进土壤呼吸的进行^[2]。Wang 等^[77]基于青藏高原草地生态系统开展了为期 7a 的增温控制试验, 发现由于降水量年际变化导致土壤呼吸随增温变化具有差异性, 而土壤水分不同是差异的主要来源。为进一步验证试验结果的准确度, Wang 等^[77]在全球尺度上收集了增温试验的数据, 证实降水显著影响土壤呼吸的增温响应, 表明增温的净效应取决于温度对土壤呼吸的促进作用与温度升高引发的土壤水分降低的抑制作用间的平衡。Guo 等^[78]基于 12a 野外增温试验也发现, 环境增温导致土壤水分下降与土壤微生物组成和功能之间具有极强的相关关系, 且土壤水分限制会削弱温度对土壤呼吸的刺激效应^[79], 因此土壤呼吸在增温环境下的变化需要综合考虑温度和水分因素。

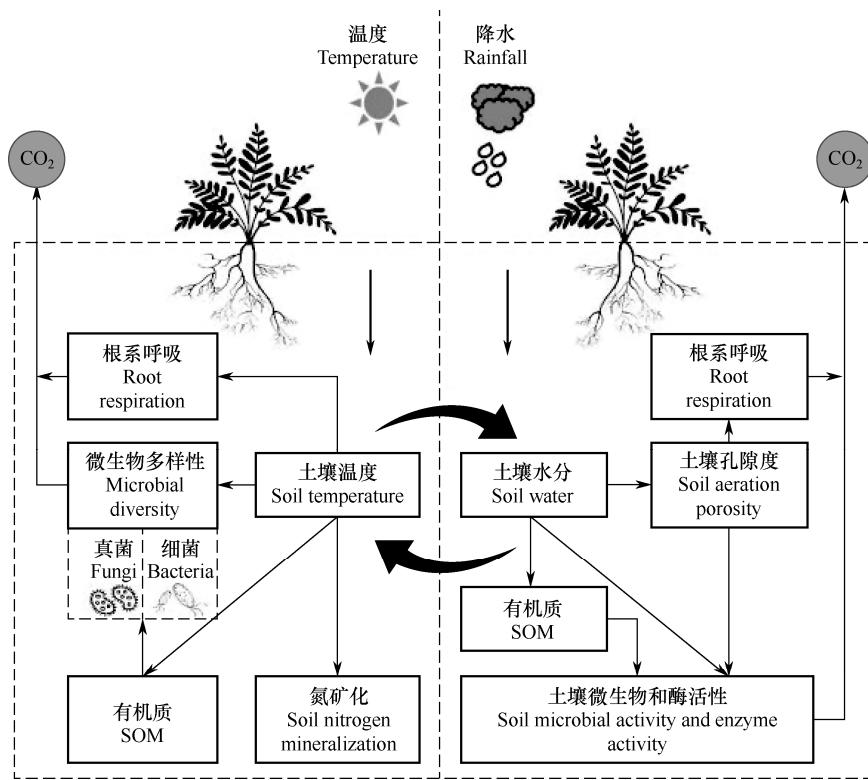


图 2 温度和水分交互作用对土壤呼吸影响示意图

Fig. 2 Schematic diagram of the effect of temperature and water interaction on soil respiration

土壤温度、水分和土壤呼吸之间的耦合关系取决于土壤水热因子的配比关系^[7], 当土壤水分成为胁迫因子时, 气候变暖引发的土壤温度升高对土壤呼吸的促进作用被干旱的负面影响所抵消; 当土壤温度成为胁迫因子时, 降水引发的土壤水分含量升高对土壤呼吸的激发效应被低温的负面影响所抑制^[2,8]。Yan 等^[80]基于山西太原天龙山自然保护区进行了为期 11a 观测试验, 发现土壤温度与水分存在明显交互作用, 采用双变量模型相较单变量模型能够更好地模拟土壤呼吸的变化状况, 因此进行土壤呼吸研究时需充分考虑土壤温度和水分的影响及其交互作用^[8]。

4 问题与展望

土壤呼吸是一个复杂的生态学过程, 温度和水分作为影响土壤呼吸最重要的两个非生物因素, 深刻理解二者对土壤呼吸的影响及其交互作用对掌握和揭示土壤呼吸规律具有重要意义。在全球尺度上尽管众多学者就土壤呼吸的生物和非生物因素开展了大量研究, 基于土壤呼吸对土壤温度和水分的响应规律已有普遍和明确的认知, 但当前气候变化是

一个长期且复杂的过程, 且鉴于土壤呼吸组分的复杂性、模拟气候变化试验存在的局限性以及不同研究的时空尺度差异性等问题, 使得土壤温湿度对土壤呼吸的影响及其调控机制尚未得到统一定论。因此, 基于陆地生态系统的土壤呼吸在未来的研究中需关注以下几个方面:

(1) 加强土壤呼吸多因素交互研究。土壤呼吸在陆地生态系统中受多种生物和非生物因素的共同影响, 而全球气候变化带来的并非某单一因素的改变, 通常是多种因素的变化同时发生, 且不同因素间存在制约和交互作用。因此, 未来应全面分析可能涉及的影响因素, 开展多因素控制试验探究不同因素对土壤呼吸的影响及其交互作用。

(2) 定量化研究土壤呼吸组分。土壤异养呼吸作为气候变暖与土壤呼吸正反馈中最不确定的因素之一, 在土壤呼吸中的占比对该反馈调节及土壤碳储量都具有重要意义。而土壤自养呼吸和自养呼吸定量化分离技术仍是现有土壤呼吸研究的技术难点, 探明植物-土壤-大气碳收支情况需在未来研究中突破土壤呼吸组分分离技术。

(3) 关注土壤呼吸对土壤初始温度和温度波动

性的响应特征。土壤初始温度是决定土壤呼吸温度敏感性的重要因子,且二者之间具有负相关关系^[81]。在平均温度相同的情况下不同温度波动对土壤微生物呼吸的影响存在差异,且温度波动和平均温度对土壤异养呼吸具有不同的调控方式,温度波动对细菌群落结构的影响大于真菌群落,并可造成土壤细菌群落和土壤酶活性之间的负反馈,造成土壤呼吸温度适应性^[82];而平均温度则是主要通过改变真菌群落及在真菌群落和酶活性之间建立正反馈关系,增强土壤呼吸对平均温度的响应。因此在进行土壤呼吸对土壤温度的响应研究时,应考虑把土壤初始温度和温度波动纳入试验设计。

(4) 探索生物多样性或群落结构组成对土壤呼吸的影响。生物多样性可提高生态系统功能,全球气候变暖导致的温室效应会对陆地生态系统生物多样性造成负面影响^[83]。土壤生物多样性在驱动陆地生态系统碳循环中起关键和主导作用^[84],因此深入探索土壤呼吸与地上/地下生物多样性及植物群落组成关联,对于全面理解陆地生态系统在全球气候变化下的碳循环具有重要意义。

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